

International School "Structure and Composition of the Lower Continental Crust"

Geophysical investigation of the LCC: 2. Results and questions A) Himalayas B) Ivrea zone

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09.10.2019.

Pavia

The underthrusted Indian lower crust: Geophysical constraints on eclogitization

György HETÉNYI

with Celso Alvizuri, Kristel Chanard, Lukas P. Baumgartner, Frédéric Herman, Rodolphe Cattin, Fabrice Brunet, Laurent Bollinger, Jérôme Vergne, John L. Nábělek, Michel Diament



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Structure and Composition of the Lower Continental Crust

09.10.2019.

The Himalaya-Tibet-... orogenic system



The Himalaya-Tibet-... orogenic system



The Himalaya-Tibet-... orogenic system





Cross-section view

Hetényi 2007 PhD Nábělek, Hetényi, Vergne et al. 2009 Science



Outline

- Introduction
- The importance of local data
- Geophysical constrains on the Indian lower crust (ILC)
- Metamorphic earthquakes



Chanard et al. in review

Local data: Moho depth

Are the rare deep earthquakes in Tibet in the mantle or the lower crust? ~26 events 1963-2001 M~4.8-6.4

IZS

Warm Tibetan

1500

upper mantle



Elev(km)

Depth (km)

0

50

100

India

shallow normal fault

Himalayan thrust

deep S. Tibet

Indian mid- to lower crust

500

- projections vs. local Moho depths
- different locations and contexts: (too) ambitious to give unique explanation for all events

ndian upper crust M

MCT

MBT

Depleted Indian mantle

1000

Distance (km)



90 km: schematic

100

2000

Local data: Moho depth

Are the rare deep earthquakes in Tibet in the mantle or the lower crust?





Local data: Geological variability

• segmentation of the Himalaya, inherited from the India plate

Hetényi et al. 2016 Sci Rep





Dasgupta et al. 1993



Local data: Dry Indian lower crust?

ARTICLE

D01: 10.1038/s41467-018-05964-1 OPEN

Lower-crustal earthquakes in southern Tibet are linked to eclogitization of dry metastable granulite

Feng Shi^{1,2}, Yanbin Wang¹, Tony Yu², Lupei Zhu³, Junfeng Zhang¹, Jianguo Wen⁴, Julien Gasc⁵, Sarah Incel⁵, Alexandre Schubnel⁵, Ziyu Li³, Tao Chen¹, Wenlong Liu¹, Vitali Prakapenka² & Zhenmin Jin¹



Xenolith samples from Tibet suggest that the Indian lower crust lacks hydrous minerals²⁷, making dehydration embrittlement less likely to operate.



- xenolith sample from >600 km to N, region without underthrusting ILC
- to date we have no relevant sample of underthrusted Indian lower crust



Geophysical constraints on the Indian lower crust

- Central Himalaya
- Seismological data
 → structure geometry
- Thermo-kinematic and petrological modelling
 → T, P, chem. → ρ
 ⇔ gravity data
- Can we constrain
 - Water content?
 - Reaction kinetics?



Approach

South

Depth (km)

Equilibrated geothern

- **Central Himalaya**
- Seismological data \rightarrow structure geometry
- Thermo-kinematic and petrological modelling \rightarrow T, P, chem. $\rightarrow \rho$ ⇔ gravity data
- Can we constrain
 - Water content?
 - **Reaction kinetics?**



CustHP feldspar Gr(HP)

Cpto(HP) feldspor On(HP) Opt(HP) q H2O

Cpx(HP) feidspie Opx(HP) 4 H2O

1000

900

4 H2O

Thermo-kinematic model Following Henry et al. 1997; Bollinger et al. 2006



Constant heat flow at the base of the model (q)

- **FEAP**: finite-element heat advection-diffusion eq.
- geometries, convergence rate are fixed
- radiogenic heat production *A*, basal heat flow *q* are varied thermal field reproduces large field datasets

Cps(HP) Pheng(HP) Gi(HP) ky coe H2O Density Cpx(HP) Pheng(HP) GITrTsl 3300 3400 2800 2900 3000 3100 3200 Gt(HP) ky law coe 2700 (kg/m^3) Cpx(HP) Pheng(HP) GITrTsPg Gt(HP) 30 ky law Cpx(HP) Pheng(HP) Gt(HP) ky q H2O Dx(HP GITrTsPg Gt(HP) 25 Cpx(HP) Cpx(HP Pheng(HP) GITrTsPg Pheng(HP) GITrTsPg H Gt(HP) Gt(HP) Cpx(HP) feldspar Gt(HP) 20 H2O Pressure (kbar) BS DX(HP) eng(HP) GITrTsPg feldspar Gt(HP) zo (GITLESPE feldspar GitHPro Bio(HP) Cpx(HP) GITrTsPg feldspar Cpx(HP) feldspar Gt(HP) Opx(HP) Gt(HP) q H2O g H2O o(HP) GITrTsPg 10 Bio(HP) GITrTsPg feldspar Gt(HP) Cpx(HP) feldspar Opx(HP) Bio(HP) GITrTsPg feldspar Opx(HP) q H2O 5 GlTrTsPg feldst Opx(HP) q 400 500 600 700 Temperature (°C) 800 900 1000 700 400 500 600 800 900

Petrogenetic grids Connolly 2005

- **Perple_X**: Gibbs energy minimization
- mineral composition is fixed: average continental lower crust

Rudnick and Fountain 1995 Rudnick and Gao 2003

1000

- solid solution phases are fixed
- water content is varied:

30

25

20

15

10

5

Pressure (kbar)

wet (all hydrous minerals), partially hydrated (amphibolitic), dry (granulitic)

Petrogenetic grids Connolly 2005



- *Perple_X*: Gibbs energy minimization
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Petrogenetic grids Connolly 2005



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Constraining model results

- temperature at key locations ⇔ A, q
- gravity anomalies:
 - far-field fit
 - ✓ partially hydrated
 - along-profile variations

misfit in the region of interest

- eclogitization kinetics:
 - 🗴 equilibrium
 - 🗴 sluggish kinetics

✓ delay



Hetényi et al. 2007 EPSL

Reaction kinetics – the role of water

250km

50km

В

Mafic 1wt.%H2O

3400

3300

3200

kg/m³)

Density

2900

2800

2700

3300

3200

3100

2800

2700

30

25

20

Α Lack of free water Overstepping of the plag-out reaction



Hetényi et al. 2007 EPSL

Update on effective reaction rate



- overstepping well-constrained
- slope ~ kinetics \rightarrow



Conclusion 1

Indian lower crust partially hydrated

• various local datasets and coupled modelling

Question 2

What are those earthquakes?

- hard to have brittle rupture at those temperatures
 - what relation to metamorphism?

Himalayan deep-crustal earthquakes

Original figures of the relocated earthquakes



Himalayan deep-crustal earthquakes

Alvizuri and Hetényi, 2019



Methodology: full-moment-tensor analysis



SO

Meticulous scrutinizing of waveform fits...

Alvizuri and Hetényi, 2019

0.14

0.97

0.64

1.71

-0.13

-2.27

98

1.01

0.05

-2.22

80

2.09

-0.37

-3.24

93

1.64

0.36

0.38

3.87

1.01

81

97



Full-moment-tensor and uncertainty

Best-fit solution away from DC towards opening crack + uncertainty

Interpretation

tectonic (DC) ∆V=0

thermal runaway $\Delta V^{\sim}0$

anticrack $\Delta V < 0$

dehydration embrittlement $\Delta V>0$

- lower crustal protolith reaches dehydration P-T conditions
- small amounts of H₂O accumulates in pores
- increased pore pressure creates fractures (smaller, then larger)
- final large fracture opens to evacuate H₂O and also slips during eqk., including/through damage (see Ben-Zion et al.)

Comparison to palpable geological examples?

• Bergen Arcs?

Austrheim et al. 1996 EPSL

- pseudotachylytes, but dry granulite \rightarrow eclogite needs external fluids
- imaginary event across Holsnøy ~ magnitude 4

Alvizuri and Hetényi, 2019

- W. Alps?
 - μ m-m documentation of dehydration–evacuation, but \rightarrow thermal runaway

Plümper et al. 2017 Nat Geosci

Comparison to laboratory experiments?

2017

Nat Comm

- Serpentinized peridotite: dehydration reactions drive stress-transfer Ferrand et al. \rightarrow acoustic emissions
- hydrated rock ③, but oceanic crust (and different experimental P) ullet

Other experiments ("but"):

- Incel et al. 2017: lawsonite-blueschist, *lws* survives the reaction and keeps H₂O ۲
- Wang et al. 2017: for mantle transition zone earthquakes ۲
- Shi et al. 2018: dry rocks, T higher by 150°C
- Incel et al. 2019: dry granulite, higher P, higher T

- Proposal for an experiment:
- continental lower crustal composition
- various amounts of H₂O

- cross at ca. 650°C / 2 GPa
- slow deformation rate

Conclusions 1 & 2

Indian lower crust partially hydrated

Metamorphic earthquakes related to dehydration

• <u>local</u> events and data – *not* a generalization for the Tibetan Plateau

"Dehydration embrittlement changes the mechanical properties of the crust, and extends the depth of the brittle rupture domain to that of the deepest hydrated phases" (Raleigh & Paterson 1965)

Deep crustal, intermediate depth, deep earthquakes

- dehydration embrittlement? transformational faulting? thermal runaway?
 - prevailing mechanism depends on P, T, chemistry & reactions, H₂O, ...

Metamorphic earthquake in the Indian lower crust beneath southernmost Tibet

Old and new geophysical results on the Ivrea body

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with Matteo Scarponi, Ludovic Baron, Théo Berthet, Jarka Plomerová, Stefano Solarino, Mattia Pistone, Luca Ziberna, Alberto Zanetti, Andrew Greenwood, Othmar Müntener

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Geological situation of the IVZ at depth

Zingg et al. 1990

Ivrea-Verbano Zone vs. Ivrea geophysical body

How deep is the Moho? ("Can we drill it?")

Gravity

First observation

Niethammer 1921

Gravity

Niggli 1946

Gravity + Seismics

Combined density and P-wave velocity models

Distance, in kilometers

Gravity

3D models

All models assume *constant* density difference!

Magnetics

First observation: \underline{V} and \underline{H}

• anomalous body continues under the sediments

Magnetics

Several 2D models Froidevaux & Guillaume 1979, Albert 1979, Lanza 1982, Belluso et al. 1990, Mouge & Galdeano 1991

First 3D model

Wagner et al. 1984

Figure 9: Sketch of the magnetic Ivrea body.

Magnetics – summary:

- IVZ top: 11 km with highs at 2.5-3 km
- in the S (Cuneo) 7-8 to 15 km
- does not (always) coincide with gravity
- no clear trend of magnetic susceptibility with rock type

Geothermics

- eighteen 2-metre probes in IVZ to obtain heat flow Haenel 1974
- rock heat production data Höhndorf 1975
- stationary thermal model Höhndorf et al. 1975

- from Curie temperature depth estimates: steady-state thermal model is questioned

Speranza et al. 2016

Seismology

Active seismics:

- difficult to get signal (energy scatters), WA fan refr.+inline refl. is best so far
- VP = 7.2-7.83 km/s reached at 3-14 km

Ansorge 1968, Berckhemer 1968, Giese 1968, Giese et al. 1973, Kissling 1984, ECORS-CROP DSS Group 1989, Hirn et al. 1989, Eva et al. 2015

Local Earthquake Tomography:

Diehl et al. 2009 VP = 7.0 km/s at 13 km, VP = 7.5 km/s at 27 km

Seismology

Local Earthquake Tomography:

- Potin 2016 PhD thesis VP = 7.0 km/s at 13 km, VP = 7.5 km/s at 18 km
- depth to "high velocity" in earlier works:
 - IVZ: 5+ km De Franco et al. 1997
 - South: 7-10 km

Solarino et al. 1997, Paul et al. 2001, Béthoux et al. 2007

• passive and active seismics comparison: gap in scales

Summary of past surveys

Method	Z to anomaly (km)	Anomaly
Gravity	0 <5 ~10	3000 kg/m ³ 3100 kg/m ³ 3300 kg/m ³
Magnetics	IVZ highs 2.5-3 IVZ elsewhere 11 (South 7-8-15)	no clear trend with lithology
Active seismics	3-14	7.2-7.83 km/s
Passive seismology	13 local highs: 5-7-10	7.0 km/s "high velocity"
Geothermics	5	(model)

All passive geophysical data/models point to top-of-anomaly at

3-10 km,

depending on location and model assumptions.

Recent campaigns

- AlpArray \rightarrow too broad scale
- IvreaArray: 10 stations @5-km spacing, 207 new gravity points

Scarponi et al. in prep.

Conclusions A few pending questions...

- Is there a pattern in lower crustal earthquakes?
- How thick is the Moho transition in active orogens?
- What is below the Bird's Head?
- Which rocks constitute the IGB?

- How could geologists and geophysicist cooperate (even) better?
- How could we better bridge across temporal and spatial scales?

